

Predicting groundwater arsenic contamination in Southeast Asia from surface parameters

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Published online: 11 July 2008; doi:10.1038/ngeo254

Arsenic contamination of groundwater resources threatens the health of millions of people worldwide, particularly in the densely populated river deltas of Southeast Asia. Although many arsenic-affected areas have been identified in recent years, a systematic evaluation of vulnerable areas remains to be carried out. Here we present maps pinpointing areas at risk of groundwater arsenic concentrations exceeding $10 \mu\text{g l}^{-1}$. These maps were produced by combining geological and surface soil parameters in a logistic regression model, calibrated with 1,756 aggregated and geo-referenced groundwater data points from the Bengal, Red River and Mekong deltas. We show that Holocene deltaic and organic-rich surface sediments are key indicators for arsenic risk areas and that the combination of surface parameters is a successful approach to predict groundwater arsenic contamination. Predictions are in good agreement with the known spatial distribution of arsenic contamination, and further indicate elevated risks in Sumatra and Myanmar, where no groundwater studies exist.

More than 100 million people worldwide ingest excessive amounts of arsenic through drinking water contaminated from natural geogenic sources. Many Asian countries, in particular, are known to be affected by high groundwater As concentrations as a result of chemically reducing aquifer conditions: Bangladesh^{1–10}, India^{3,11,12}, China^{13,14}, Nepal¹⁵, Cambodia^{16–18} and Vietnam^{17,19,20}. However, because As analysis is expensive and time consuming, groundwater resources of many regions still remain to be tested. Therefore, maps pinpointing areas vulnerable to As contamination can guide households at risk of arsenic contamination, as well as scientists and policy-makers, to initiate early mitigation measures and protect the populations from chronic As poisoning.

Though the exact chemical conditions and reactions leading to As mobilization are still under debate, it is generally accepted that microbial and/or chemical reductive dissolution of As-bearing iron minerals in the aquifer sediments^{1,4,21} is the main cause for the release of As. Reducing conditions are often associated with the presence of natural (bio)degradable organic carbon embedded in sediments^{9,11,22–25}. Other identified key characteristics of contaminated areas are rapidly buried, young (Holocene) sediments and low hydraulic gradients in flat and low-lying areas^{8,18,26,27}. Ideally, an As prediction model for groundwater should be based on parameters that indicate the key characteristics mentioned above in three dimensions. However, in the absence of a three-dimensional spatially continuous database of aquifer conditions to depth, globally and regionally available (two-dimensional) surface parameters can be used as indicators for As enrichment in underlying aquifers^{28,29}.

In the past, several geostatistical interpolation methods (for example kriging) have been used to predict elevated As in groundwater on a regional scale^{30–32}. However, for predictions of areas where no groundwater quality data exist, interpolation

methods are not applicable³² and models based on logistic regression are more appropriate^{33,34}. In an expert-based statistical model to delineate areas at risk of groundwater As contamination on a coarse global scale, we found that geological information was of crucial importance²⁹. Here we focus on an in-depth assessment of depositional environments in Southeast Asia. We use a logistic regression approach based on relationships between sedimentary information, soil maps and measured groundwater As data of Bangladesh, Cambodia and Vietnam to assess the relative importance of the different surface proxies in these countries. We apply these relationships to set up prediction maps of As contamination in Southeast Asia including Indonesia (Sumatra) and other countries where groundwater quality data are scarce (Myanmar and Thailand). Furthermore, we verify the predicted risk in South Sumatra, where the groundwater has not previously been tested for the presence of As.

ARSENIC PREDICTION MODEL

Our model is based on three assumptions. First, sedimentary depositional environments are characterized by a unique combination of chemical, physical and biological properties³⁵ and can serve as indicators (proxies) for chemical and physical conditions of the aquifers beneath the surface. Second, soil properties are proxies for present and past drainage conditions and they are also indicators of recent depositional environments. Third, soil textures, for example clay and silt, are proxies for the chemical maturity of the sediments, where clay is more mature than silt. An important factor in the development of soil textures is topography³⁶, which enables the delineation of areas where the model is applicable.

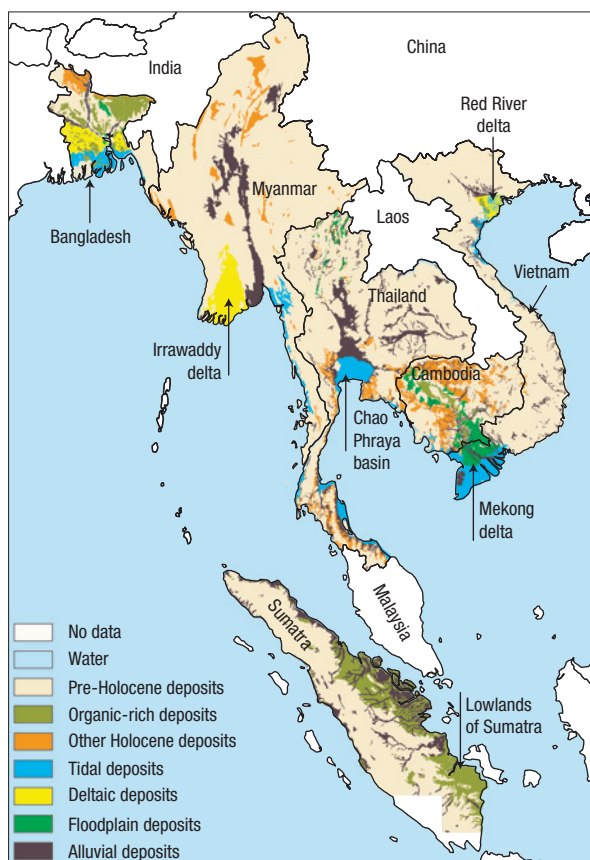


Figure 1 Uniformly classified geological map of Southeast Asia. Seven different sedimentary depositional environments are indicated in the mapped countries of Bangladesh, Myanmar, Thailand, Cambodia, Vietnam and Sumatra (Indonesia).

Geographic information system datasets were established from digital elevation data, countrywide geological maps and global soil data (Food and Agriculture Organization of the United Nations), which were converted to a raster format using ArcGIS (Version 9.2). An overview of geographic information system data used in this study is provided in the Supplementary Information, Table S1. Because each geological map applied a different classification terminology, we created a uniformly classified geological map for all regions (Fig. 1). Although Bangladesh does not geographically belong to Southeast Asia, it was included in the model because of the large number of available data points⁵. Statistical relations between As concentrations and 30 parameters related to soil properties, geology, climate and hydrology (see Supplementary Information, Table S2) were initially evaluated by stepwise regression. The six parameters showing a significance greater than 95% and two additional soil parameters were used in the final model (see the Methods section and Table 1). Because young geological deposits and As groundwater contamination are rarely observed in areas with steeper slopes, and groundwater As concentration data are available only for regions with a flat topography (slope $< 0.1^\circ \approx 0.17\%$), areas with slopes greater than 0.1° were excluded from the model (see Supplementary Information, Fig. S1).

The prediction required three steps: (1) aggregation and binary-coding of measured As concentrations to diminish spatial heterogeneities (dependent variables), (2) logistic regression to

Table 1 Results of logistic regression analysis. Weighting coefficients of the independent variables (λ) were used to calculate probabilities of As contamination. Wald values (%) indicate the relative importance of the variables, and p -values the absolute significance, where a value less than 0.05 indicates a significance of at least 95%. Variables that were not statistically significant (on the basis of a 95% level) were excluded from the model (see Supplementary Information, Table S2), with the exception of medium-textured soils and silt in subsoil (see the Methods section). Excluded variables are tidal deposits, other Holocene deposits, pre-Holocene deposits (Fig. 1), coarse-textured soils, soil sand and clay contents, and climate.

Variables		λ	Wald	p -value
Sedimentary depositional environments	Deltaic deposits	1.65	71.55	<0.001
	Organic-rich deposits	1.11	72.20	<0.001
	Alluvial deposits	0.55	12.51	<0.001
	Floodplain deposits	-0.95	2.32	0.010
Soil variables	Medium-textured soils	0.60	7.80	0.128
	Fine-textured soils	0.24	1.00	0.005
	Silt in subsoil	0.10	6.62	0.317
	Silt in topsoil	-0.09	6.28	0.012
—	Regression constant	-1.54	20.67	<0.001

obtain weighting coefficients of independent variables (see the Methods section and corresponding results in Table 1) and (3) calculation of the probability of As contamination on the basis of the threshold value of $10 \mu\text{g l}^{-1}$. The spatial datasets considered as independent variables for the model are topographic data to delineate the model area, sedimentary depositional environments as a proxy for aquifer conditions and soil variables as a proxy for drainage and chemical maturity of sediments (see the Methods section).

SURFACE PARAMETERS CONTRIBUTING TO THE MODEL

The weighting factors (λ) and significance of the eight independent variables retained in the final model are listed in Table 1. In general, variables describing sedimentary depositional environments have a larger contribution to the model than soil variables, presumably because geological variables come closest to describing As contamination in the aquifer itself. The presence of young deltaic deposits ($\lambda = 1.65$) is a particularly significant indicator for As-contaminated aquifers. In Southeast Asia, delta initiation and progradation occurred simultaneously with the Holocene Climate Optimum³⁷. Therefore, delta progradation resulted in the burial of organic matter at a high rate. The presence of relatively fresh organic carbon provides favourable conditions to establish reducing environments, which may lead to As enrichment in groundwater. Logistic regression confirms that organic-rich deposits ($\lambda = 1.11$) play an important part in the model. Recent alluvial deposits ($\lambda = 0.55$) are also indicative of elevated As concentrations in groundwaters (Table 1). Of the soil parameters, medium-textured soils seem to be indicative of As-bearing aquifers evolved from rapid accumulation of young (Holocene) sediments. In contrast, the negative weighting coefficient ($\lambda = -0.95$) for floodplain deposits implies that these fine-grained deposits could overlie aquifers low in dissolved As (compare Fig. 1 and Fig. 3 below). Fine-grained deposits (high clay content) thereby point to low-energy depositional environments with condensed sediments where Holocene aquifers are rare and where groundwater is probably drawn from older aquifers.

Pre-Holocene deposits, other Holocene deposits and tidal deposits (Fig. 1) were found to be statistically insignificant ($p > 0.05$) and they were excluded from the model in the first

stepwise regression (see Supplementary Information, Table S2 and Table 1). Tidal sediments are generally associated with aquifers abundant in sulphate, which may be microbially reduced to sulphide and re-precipitate As (refs 38,39). They might serve as proxies for low-risk areas, but our As data in such aquifers were possibly too few to see such a relation.

INFERRING DEPOSITIONAL ENVIRONMENTS AT DEPTH FROM THE SURFACE

As mentioned in the introduction, our model is based on two-dimensional data (that is, surface maps). Nevertheless, geological information (sedimentary depositional environments) inherently contains a three-dimensional component. The recent sedimentary history of major Southeast Asian basins is characterized by delta initiation, which occurred on a global scale at about 8,500–6,500 years BP and was principally controlled by the deceleration of sea-level rise⁴⁰. Delta progradation resulted in the unconformable deposition of thick late Pleistocene–Holocene sediments on Pleistocene and older sediments, for example as incised-valley-fill deposits^{8,18}. As-contaminated aquifers are mainly present in these Holocene aquifers, whereas deeper Pliocene–Pleistocene aquifers are, to a large extent, free of As (refs 7,8,26). The boundary between Pleistocene and Holocene sediments is not located at a constant depth⁸, and this can lead to misclassifications, because our model inherently assumes that the underlying aquifer belongs to the same sedimentary depositional environment as the surface.

Two situations exist where the environment at the surface does not reflect the geology at depth, (1) when the As measurements used in the model were obtained from tube-wells tapping deep (Pleistocene) aquifers because shallow (Holocene) aquifers are not present, or shallow groundwater is too saline for consumption (for example in coastal regions), and (2) when Holocene sediments were deposited at sedimentation rates too small to form a usable aquifer. In both situations Holocene depositional environments are present at the surface and high-risk areas are indicated, although measured As concentrations are low (false-positive cases). Even though the probability maps would be improved by including the thickness of Holocene sediments, the absence of country-wide three-dimensional geological data rules this out. Furthermore, the complexity of aquifer heterogeneity at a local scale makes it inevitable that misclassifications occur.

PROBABILITY MAPS

Predicted areas at risk of As contamination agree well with known spatial contamination patterns, a finding that is supported by the results of the model classification indicating the performance of model prediction (Fig. 2) and the Hosmer–Lemeshow goodness-of-fit statistics (see Supplementary Information, Table S3). An absolute average deviation of 7.3% is found between expected and modelled probabilities of As being less than or equal to $10\ \mu\text{g l}^{-1}$ or more than $10\ \mu\text{g l}^{-1}$. The model is further characterized by the receiver operating characteristic curve area of ~ 0.7 (see Supplementary Information, Fig. S2), which is a good result considering that neither depths of analysed groundwater wells nor aquifer hydrological data form part of the model.

The probability maps for As concentrations exceeding $10\ \mu\text{g l}^{-1}$ are presented in Fig. 3 and Supplementary Information, Fig. S3 (probability map of the whole of Southeast Asia). The highest probabilities (0.7–0.8) for As contamination ($>10\ \mu\text{g l}^{-1}$) are found in the south-central part of Bangladesh, with a value of 0.5–0.6 in the northeastern Sylhet basin (Fig. 3a). The probability of finding contaminated wells in the Red River delta reaches a value of 0.7 (Fig. 3e). The sedimentary depositional

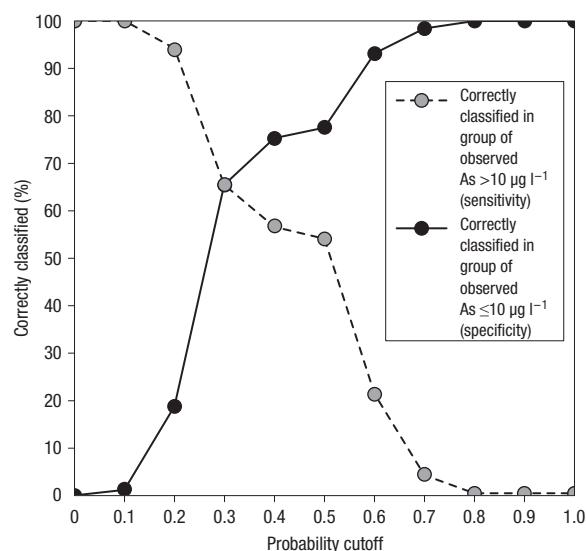


Figure 2 Model classification results. The graph shows the sensitivity (true positives) and specificity (true negatives) of the model for different probability cutoff values. The full classification table is provided in the Supplementary Information, Table S4. A probability threshold of 0.4 was applied to delineate low- and high-risk areas in the binary risk maps shown in Figs 3b,d and 4b (see the Methods section).

environments present along the Mekong river differ from the deltaic environments of Bangladesh and the Red River delta in that organic-rich deposits are found at close distance to the modern Mekong and Bassac river courses, surrounded by extensive floodplain areas free of As (ref. 18) (compare Figs 1 and 3). The probability map for the Mekong delta shows values of up to 0.6 at close distance to the modern Mekong and Bassac river courses and adjacent swampy marshes (Fig. 3c). In addition, the large floodplain of Lake Tonle Sap, with organic-rich sediments, is a risk area, with probabilities ranging between 0.4 and 0.6.

To compare our predicted low-risk (probability ≤ 0.4) and high-risk areas (probability > 0.4) (Fig. 3b,d) with measured As contamination, a misclassification analysis was carried out on the basis of 1,756 aggregated and binary-coded (≤ 10 or $> 10\ \mu\text{g l}^{-1}$) As measurement data for the three deltas. We report 71% (1,023 aggregated points), 59% (107 a.p.) and 75% (100 a.p.) of correctly classified cases for Bangladesh, Red River and Mekong deltas respectively (average 70%). In our study, the Red River and Mekong deltas have 6 and 9% false-negative cases. In Bangladesh, false-negative cases (13%) were found specifically in wells lying at a close distance to rivers. The number of false negatives is outnumbered by false positives in all three deltas, with 17, 35 and 16% for Bangladesh, the Red River and Mekong deltas, respectively. Errors in the prediction can arise from the commonly reported well-to-well variability, where wells with low As levels are often present at close distance from wells high in arsenic^{5,6,18,19}, as well as from uncertainties in measured As (estimated at 25%) leading to misclassifications of concentrations close to the threshold of $10\ \mu\text{g l}^{-1}$. However, we interpret these misclassifications to be mainly an effect of modelling three-dimensional processes on the basis of two-dimensional data.

Apart from the three deltas discussed above, our Southeast Asia probability map (see Supplementary Information, Fig. S3) also highlights risk areas that are largely unknown or unreported, particularly in Sumatra, Myanmar and Thailand (Fig. 3f).

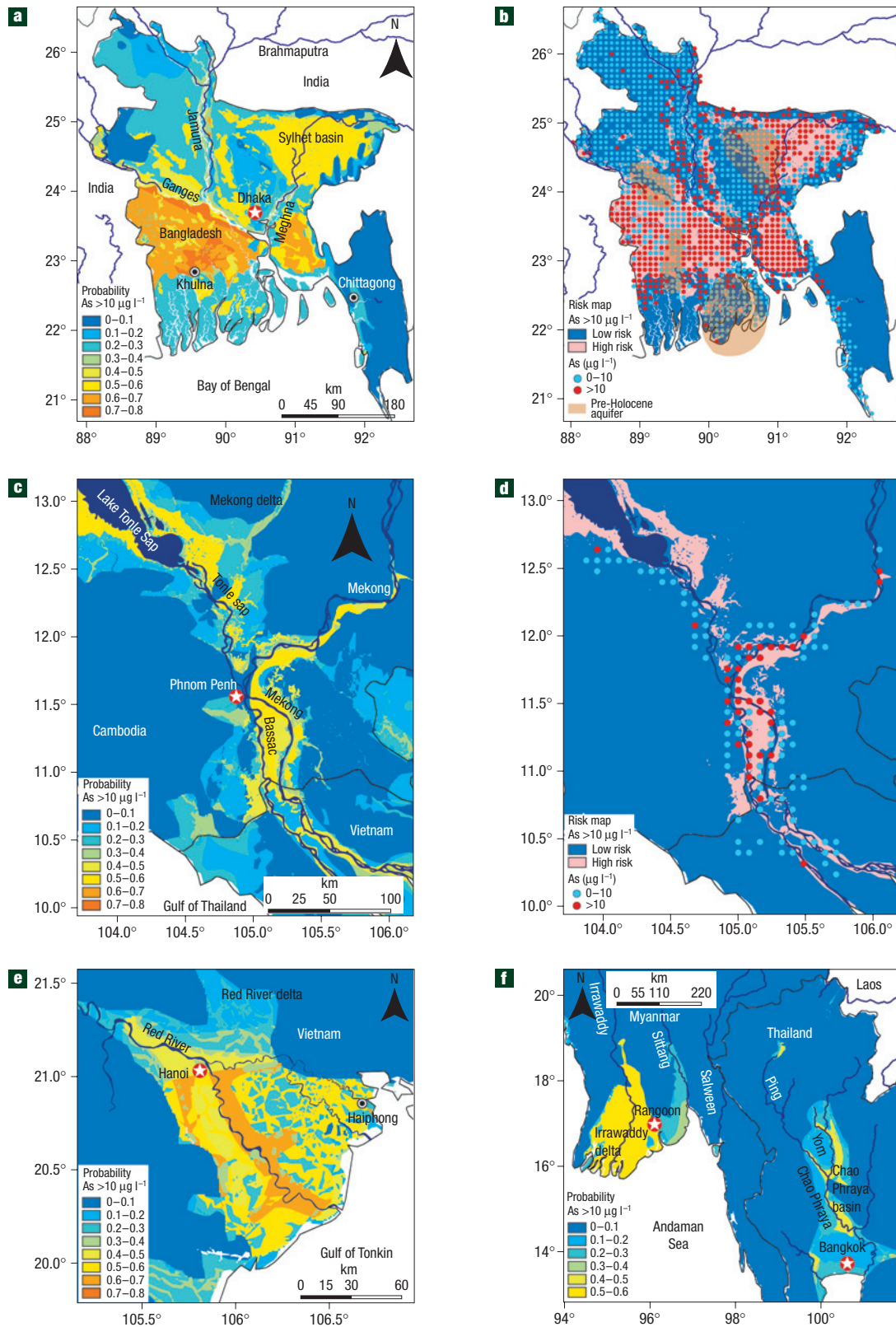


Figure 3 Modelled probability of As concentrations exceeding $10 \mu\text{g l}^{-1}$ under reducing aquifer conditions. **a**, Continuous probability map of Bangladesh. **b**, Binary map of Bangladesh (probability threshold 0.4) indicating high- and low-risk areas overlain by aggregated As concentrations. Areas where groundwater is mainly drawn from Pleistocene aquifers are sketched in brown. **c**, Continuous probability map of the Mekong delta (Cambodia and Vietnam). **d**, Binary risk map of the Mekong delta overlain by aggregated As concentrations. **e**, Continuous probability map of the Red River delta (Vietnam). **f**, Continuous probability map of the Irrawaddy delta (Myanmar) and Chao Phraya basin (Thailand).

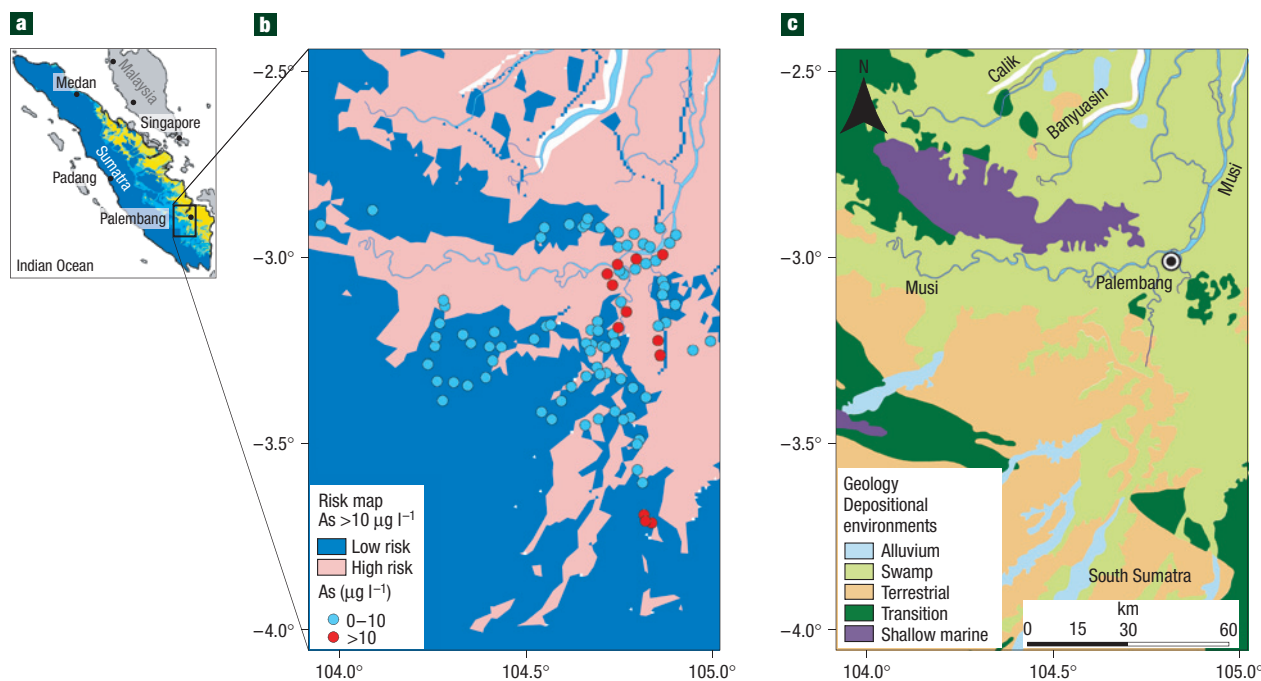


Figure 4 Maps of the model verification study area in Southeast Sumatra. **a**, Map of the whole of Sumatra depicting the probability of groundwater As concentrations exceeding $10 \mu\text{g l}^{-1}$. The corresponding colour code is given in Fig. 3. **b**, Binary risk map (probability cutoff 0.4) and As concentrations measured for this study in the vicinity of Palembang (South Sumatra). Swampy areas (high-risk area) are scarcely populated and the number of sampled tube-wells in this area is therefore limited. **c**, Geological map (source: see Supplementary Information, Table S1).

VERIFICATION OF PREDICTIONS FOR SUMATRA

According to the modelled probability, an area of about $100,000 \text{ km}^2$ on Sumatra's east coast is prone to high risk of As contamination (probability > 0.4) (see Fig. 4a and Supplementary Information, Fig. S3). To validate the Sumatra prediction map, 97 groundwater samples were collected in 2007 in the Province of South Sumatra in the vicinity of Palembang (see the Methods section). This area was chosen because it is at the border of a low- and high-risk area and not previously studied. Because there is no large variation in geological and topographical features along Sumatra's east coast, the study area is representative of the whole high risk area. Figure 4b shows As concentrations ($\leq 10 \mu\text{g l}^{-1}$ and $> 10 \mu\text{g l}^{-1}$) measured in the groundwater, imposed on the binary probability map. The classification results for Sumatra (63% correctly classified, 36% false positive and 1% false negative) are comparable to those of the Bangladesh, Red River and Mekong deltas. In total, 94% of the 50 tube-wells located in the low-risk area have As levels below $10 \mu\text{g l}^{-1}$. Of the 12 contaminated wells (As $> 10 \mu\text{g l}^{-1}$) 75% are positioned in the high-risk area. However, in the high-risk area the contaminated wells are clearly outnumbered by uncontaminated groundwater measurements (nine contaminated wells versus 38 uncontaminated wells), for reasons explained below.

On average, both high- and low-risk areas in Sumatra are characterized by high dissolved organic carbon, NH_4^+ , HCO_3^- , PO_4^{3-} and Fe(II) and low SO_4^{2-} concentrations (Supplementary Information, Table S5). At least two-thirds of the sampled groundwaters are reducing in nature, especially those located in the high-risk area. Because the chemical conditions of the aquifers would permit the reductive dissolution of As, other explanations for the overall low As concentrations must be considered.

To gain a greater insight into the characteristics of the aquifer, the local geology was examined laterally and in depth. The studied area is geologically young (Tertiary–Quaternary) (Fig. 4c). The sediment contains deposits from peat-swamp forests that developed during the Holocene era (past 5,000 to 10,000 years) and unconformably overlie older sediments⁴¹. The outline of the high-risk As-contamination zone mainly follows the outline of these deposits. The peat deposits are usually found ranging from 4 m to 8 m, although depths of up to 24 m have been reported⁴². However, the depths of sampled tube-wells in Sumatra have a median of 40 m, which implies that most of them tap groundwaters from aquifers below the Holocene peat deposits. This shows that the prediction map is a useful tool for the identification of areas at risk of As contamination, but that understanding the local geology as a function of depth is of vital importance for specific areas.

ARSENIC CONTAMINATION IN THAILAND AND MYANMAR

The probability map in Fig. 3f shows that the Chao Phraya basin in central Thailand and the Irrawaddy delta in Myanmar have a risk of elevated As being present in groundwater, whereas the Sittang basin (Myanmar) has a lower risk, and the Salween basin (Myanmar) has virtually no risk at all. The problem of As contamination in the Irrawaddy delta is partly known⁴³, although its spatial extent has not been investigated to date, which is particularly worrying considering the size of the area at risk. In the Chao Phraya basin in central Thailand, a groundwater survey was undertaken in 2001 using As field test kits to test wells with minimum depths of 80 m (ref. 44). The range of measured As concentrations in the 37 tested wells was $< 1\text{--}100 \mu\text{g l}^{-1}$ with an average of $11 \mu\text{g l}^{-1}$. These results

correlate with our maps, which predict a low to moderate As contamination. Indeed, north–south geological profiles across the basin⁴⁵ indicate maximal depths of 20 m for the pre-Holocene incisions, implying that sampled tube-wells draw water from Pleistocene or older aquifers.

In contrast to the slow sedimentation rates of the Sittang and Chao Phraya deltas, the Irrawaddy delta received massive amounts of sediments during the Cenozoic era⁴⁶. The Irrawaddy River⁴⁷ still has a ten times larger sediment load than the Chao Phraya River⁴⁵. In 2002 the Departments of Medical Research and Health (Lower Myanmar), financially supported by UNICEF, conducted a groundwater sampling campaign in the Irrawaddy division⁴³. In total, 99 groundwater samples (90 shallow tube-wells and nine dug wells) were collected in 25 villages, and As was quantified by atomic absorption spectroscopy. It was reported that 67% of the sampled wells had As levels greater than 50 µg l⁻¹. These results show that the risk of As contamination in the Irrawaddy delta, as indicated by the probability map, should be taken seriously, and that there is an urgent need to test shallow tube-wells in other townships or districts in the high-risk area.

ASSESSING ARSENIC ELSEWHERE

In the study presented here, we identified regions in Southeast Asia, on the basis of As prediction maps, where tube-wells should be tested for elevated As concentrations (>10 µg l⁻¹). Although the use of such maps in other scientific investigations (for example climate research) is a common practice, the prediction of As (and other groundwater conditions) is still an emerging technique in the field of natural groundwater contaminants. Our As model differs from earlier models in its ability to predict contamination in areas of unknown groundwater quality and on a subcontinental scale. The strength of the prediction lies rather in the combination of surface parameters than in the individual parameters. As an example, the Sittang basin and the vicinity of the lower Mekong River branches are both characterized by alluvial deposits, but only the latter has a modelled high risk because of the contribution of soil properties. A limitation to be considered is that shallow, As-bearing groundwater can be expected only where Holocene aquifers are present. Where this is not the case, high-risk areas may indicate the presence of reducing aquifers, but As concentrations in groundwater could be lower than predicted.

Our approach provides a blueprint for further modelling and mapping of As-tainted aquifers in and outside Southeast Asia. The probability maps can further be improved when data with higher spatial resolution or in three dimensions become available, although we emphasize that it will not be possible to account for the local heterogeneities of aquifers. The presented prediction maps are a valuable and resource-saving tool that can serve both scientists and policy-makers to initiate early mitigation measures to protect the people from As-related health problems as well as to efficiently guide water resource management.

METHODS

Geological maps of Bangladesh, Cambodia, Thailand and Vietnam were available in digital format (see Supplementary Information, Table S1). Maps of Myanmar and Sumatra were digitized for this study. Geological maps of Malaysia and Laos were not available. The sedimentary depositional environments used in the model as independent variables are deltaic deposits, organic-rich sediments (for example sediments deposited in marshy environments), floodplains, alluvial deposits, tidal deposits, other Holocene sediments and pre-Holocene sediments. Soil variables are percentages of silt, clay and sand in both the topsoil (0–30 cm) and subsoil (30–100 cm) and coarse, medium and fine soil textures.

We compiled more than 4,600 data points of groundwater As concentrations from Bangladesh (median well depth 35 m), the Mekong delta (Cambodia and Southern Vietnam, m.w.d. 39 m) and the Red River delta (Northern Vietnam, m.w.d. 30 m). The data originate from BGS and DPHE (Bangladesh, $n = 3,534$) (ref. 5), from Buschmann *et al.* (Mekong delta, $n = 352$) (refs 18,20) and from Berg *et al.* (Red River delta, $n = 720$) (refs 17,25,48) and were used without a restriction in well depths. To test the model, 97 tube-wells were randomly sampled for this study in Sumatra at about 5–10 km intervals at an average sampling density of one sample per 54 km². This study area is positioned at a latitude of 2° 872 S to 3° 911 S and a longitude of 103° 949 E to 104° 993 E (sampling area 85 km by 65 km). Procedures of sampling and analysis were carried out as described in ref. 18. Concentrations of As and additional parameters measured in Sumatra groundwater samples are provided in Supplementary Information, Table S5.

Arsenic concentrations are point measurements within vertical depths of the wells, whereas the other variables have coarser spatial resolution, generally greater than 30 arc seconds. Point data of measured As concentrations were therefore aggregated using the geometric mean to a resolution of one point per pixel with a size of 5 arc minutes (~9.3 km at the equator), which is the pixel size of the global soil data (Food and Agriculture Organization of the United Nations). Aggregation resulted in a decreased dataset of 1,756 pixel-based data points (Bangladesh 1443, Red River delta 180 and Mekong delta 133 points). The aggregated point-data were binary-coded using the World Health Organization guideline value for As in drinking water (10 µg l⁻¹) as a threshold. We acknowledge that several countries apply a guideline of 50 µg l⁻¹, but adopting this threshold would result in significantly fewer data points (992) for the calibration of the model. The binary variable of whether As concentrations exceed the World Health Organization threshold (1) or not (0) was applied as the dependent variable in this study.

Logistic regression was used to determine the weighting of the independent variables. This is a common statistical method in environmental research and enables the concurrent use of continuous and categorical variables^{33,34}. The parameter P denotes the probability of As concentrations exceeding the World Health Organization threshold. Logistic regression models $\log(\text{odds})$, which is defined as the ratio of the probability that an event occurs to the probability that it fails to occur, $\log(P/(1-P))$, as a linear combination of independent variables⁴⁹,

$$\log(\text{odds}) = C + \sum_{i=1}^n \lambda_i X_i,$$

where C is the intercept of regression, X_i are independent variables and λ_i are the weighting coefficients that were obtained using the maximum-likelihood procedure⁴⁹. Exponential values of coefficients, Wald statistics and p -values (Table 1) indicate the importance of each variable. Statistically insignificant independent variables were excluded from the model during each of the subsequent regression steps (Supplementary Information, Table S2). The threshold for maintaining a variable in the model was determined by the 95% significance level ($p < 0.05$). The silt contents in the subsoil and medium-textured soils were kept in the model because of their good spatial match with known contaminated areas, which is supported by the presence of silty sands at the surface of regions showing contaminated aquifers in West Bengal⁵⁰, and by reported elevated groundwater As concentrations in aquifers capped with fine surface material (silt and clay) in Bangladesh^{10,27}.

According to the calculated odds, the probability (P) of there being an As concentration above 10 µg l⁻¹ was calculated as follows:

$$P = \frac{\exp(C + \sum_{i=1}^n \lambda_i X_i)}{1 + \exp(C + \sum_{i=1}^n \lambda_i X_i)}.$$

In addition, we tested how successfully the model predicted the number of contaminated cases for probability intervals of 0.1 (Fig. 2 and Supplementary Information).

Misclassification occurs when either a point with As concentration greater than 10 µg l⁻¹ falls in an uncontaminated area (false negative) or an As concentration less than 10 µg l⁻¹ falls in a contaminated area (false positive). On the basis of the model classification results (Fig. 2 and Supplementary Information, Table S4), a probability threshold of 0.4 resulted in a minimum misclassification rate. The risk maps were hence categorized into low-risk areas (probability < 0.4) and high-risk areas (probability > 0.4).

Received 21 January 2008; accepted 17 June 2008; published 11 July 2008.

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Supplementary Information accompanies this paper on www.nature.com/naturegeoscience.

Acknowledgements

We thank T. Rosenberg, R. Febriamansyah, M.A. Hayatuddin and E. Nofyan for support during groundwater sampling in Sumatra; C. Stengel, T. Rüttimann, M. Langmeier and R. Illi for elemental analyses; K. Abbaspour and B. den Brok for discussions; R. Wildman and H. Rowland for proofreading and the anonymous reviewers for comments.

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